# PRACTICAL pH MEASUREMENTS ON NATURAL WATERS

# A. K. Covington and W. Davison

(1) Dilute solutions and freshwater including 'acid-rain' samples  $(I < 0.02 \text{ mol kg}^{-1})$ 

Major problems could be encountered due to errors associated with the liquid junction. It is recommended that either a free diffusion junction is used or it is verified that the junction is working correctly using dilute solutions as follows. For commercial electrodes calibrated with IUPAC aqueous RVS or PS standards, the pH(X) of dilute solutions should be within ±0.02 of those given in Table 1. The difference in determined pH(X) between a stirred and unstirred dilute solution should be < 0.02. The characteristics of glass electrodes are such that below pH 5 the readings should be stable within 2 min, but for pH 5 to 8.8 or so minutes may be necessary to attain stability. Interpretation of pH(X) measured in this way in terms of activity of hydrogen ion,  $a_{H_+}$  is subject<sup>1</sup> to an uncertainty of ±0.02 in pH.

#### (2) Seawater

Measurements made by calibration of electrodes with IUPAC aqueous RVS or PS standards to obtain pH(X) are perfectly valid. However, the interpretation of pH(X) in terms of the activity of hydrogen ion is complicated by the non zero residual liquid junction potential as well as by systematic differences between electrode pairs, principally attributable to the reference electrode. For 35‰ salinity seawater (S = 0.035)  $a_{H_+}$  calculated from pH(X) is typically 12% too low. Special seawater pH scales have been devised to overcome this problem:

(i) The total hydrogen ion scale,  $pH_{\eta}$ , is defined in terms of the sum of free and complexed (total) hydrogen ion concentrations, where

$${}^{T}C_{H} = [H^{+}] + [HSO_{4}] + [HF]$$
  
So,  $pH_{T} = -\log {}^{T}C_{H}$ 

Calibration of the electrodes with a buffer having a composition similar to that of seawater, to which  $pH_T$  has been assigned, results in values of pHT(X) (Tables 2, 3) which are accurately interpretable in terms of  ${}^{T}C_{H}$ .

(ii) The free hydrogen ion scale,  $pH_{p}$  is defined, and fully interpretable, in terms of the concentration of free hydrogen ions.

$$pH_{F} = -\log[H^{+}]$$

Values of  $pH_{\rm F}$  as a function of temperature have been assigned to the same set of  $pH_{\rm T}$  seawater buffers,and so alternatively can be used for calibration (Tables 2, 3) <sup>2,3</sup>

### (3) Estuarine water

Prescriptions for seawater scale buffers are available for a range of salinities. Reliable estuarine pH measurements can be made by calibrating with a buffer of the same salinity as the sample. However, these buffers are difficult to prepare and their use presumes prior knowledge of salinity of the sample. Interpretable measurements of estuarine pH can be made by calibration with IUPAC aqueous RVS or PS standards if the electrode pair is additionally calibrated using a 20‰ salinity seawater buffer.<sup>4</sup> The difference between the assigned pH<sub>SWS</sub> of the seawater buffer and its measured pH(X) value using RVS or PS standards is

$$\Delta pH = pH_{sws} - pH(X)$$

Values of  $\Delta pH$  should be in the range of 0.08 to 0.18. It empirically corrects for differences between the two pH scales and for measurement errors associated with the electrode pair. The pH(X) of samples measured using IUPAC aqueous buffers, can be converted to pH<sub>T</sub> or pH<sub>F</sub> using the appropriate measured  $\Delta pH$ :

$$pH_T = pH(X) - \Delta pH$$
  
or  $pH_F = pH(X) - \Delta pH$ 

This simple procedure is appropriate to pH measurement at salinities from 2‰ to 35‰. For salinities lower than 2‰ the procedures for freshwaters should be adopted.

## References

- 1. Davison, W. and Harbinson, T. R., Analyst, 113, 709, 1988.
- Culberson, C. H., in *Marine Electrochemistry*, Whitfield, M. and Jagner, D., Eds., Wiley, 1981.
- 3. Millero, F. J., Limnol. Oceanogr., 31, 839, 1986.
- Covington, A. K., Whalley, P. D., Davison, W., and Whitfield, M., in *The Determination of Trace Metals in Natural Waters*, West, T. S. and Nurnberg, H. W., Eds., Blackwell, Oxford, 1988.
- 5. Koch, W. F., Marinenko, G., and Paule, R. C., *J. Res. NBS*, 91, 33, 1986.

	Ionic strength mmol kg <sup>-1</sup>	Concentration(x) mmol kg <sup>-1</sup>	pH $p_{CO_2} = 0$	pH p <sub>CO2</sub> = air
Potassium hydrogen phthalate	10.7	10	4.12	4.12
	1.1	1	4.33	4.33
xKH <sub>2</sub> PO <sub>4</sub> + $x$ Na <sub>2</sub> HPO <sub>4</sub>	9.9	2.5	7.07	7.05
$x \text{KH}_2 \text{PO}_4 + 3.5 x \text{Na}_2 \text{HPO}_4$	10	0.87	7.61	7.58
$Na_2B_4O_7 \cdot 10H_2O$	10	5	9.20	_
HCl	0.1	0.1	4.03	4.03
SRM2694-I <sup>a</sup>	_	_	4.30	_
SRM2694-IIª	_	_	3.59	_

### TABLE 1. pH of Dilute Solutions at 25°C, Degassed and Equilibrated with Air, Suitable as Quality Control Standards

*Note:* The pH of solutions near to pH 4 is virtually independent of temperature over the range of 5 to 30°C.

<sup>a</sup> Simulated rainwater samples are available (Reference 5) from NIST containing sulfate, nitrate, chloride, fluoride, sodium, potassium, calcium and magnesium.

TABLE 2. Composition of Seawater Buffer of Salinity S = 35‰ at 25°C (Reference 3)							
Solute	mol dm <sup>-3</sup>	mol kg <sup>-1</sup>	$\mathbf{g}  \mathbf{k} \mathbf{g}^{-1}$	g dm⁻³			
NaCl	0.3666	0.3493	20.416	20.946			
Na <sub>2</sub> SO <sub>4</sub>	0.02926	0.02788	3.96	4.063			
KCl	0.01058	0.01008	0.752	0.772			
CaCl <sub>2</sub>	0.01077	0.01026	1.139	1.169			
MgCl <sub>2</sub>	0.05518	0.05258	5.006	5.139			
Tris	0.06	0.05717	6.926	7.106			
Tris · HCl	0.06	0.05717	9.010	9.244			

Tris = tris(hydroxymethyl)aminomethane (HOCH<sub>2</sub>)<sub>3</sub>CNH<sub>2</sub>.

A 20‰ buffer is made by diluting the 35‰ in the ratio 20:35.

### TABLE 3. Assigned Values of 20‰ and 35‰ Buffers on Free and Total Hydrogen Ion Scales. Calculated from Equations Provided by Millero (Reference 3)

	pН <sub>т</sub>	рН <sub>т</sub>	pH <sub>F</sub>	pН <sub>г</sub>
Temp (°C)	S = 20%	<i>S</i> = 35‰	S = 20‰	S = 35‰
5	8.683	8.718	8.759	8.81
10	8.513	8.542	8.597	8.647
15	8.351	8.374	8.442	8.491
20	8.195	8.212	8.292	8.341
25	8.045	8.057	8.149	8.197
30	7.901	7.908	8.011	8.059
35	7.762	7.764	7.879	7.926